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Drivers of temperature sensitivity of decomposition of soil organic matter along a mountain altitudinal gradient in the Western Carpathians

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Abstract Mountain forest soils contain an important stock of carbon. Their altitudinal gradient can serve as a model for research on the potential risk of increased emission of carbon dioxide to the atmosphere, in a positive feedback of global warming. Using soil samples collected at three elevations (600, 900, and 1200 m a.s.l.) from five separate slopes of the Carpathian Mountains (Poland), we studied the effects of soil physical, chemical and microbial properties controlling the temperature sensitivity (Q_{10} values) of organic matter decomposition in forest soils. Data of soil basal respiration rate measured in laboratory conditions at six different temperatures (5, 10, 15, 20, 25 and 30 °C) were fitted to a Gaussian function. The modelled soil respiration rates differed between altitudes at temperature exceeding 15 °C, and the respiration rate of soil from 1200 m a.s.l. was higher than in soils from the two lower elevations. Based on the modelled respiration values, we calculated Q_{10} values in the low (Q_{10L} , 0–10 °C), medium (Q_{10M} , 10–20 °C) and high (Q_{10H} , 20–30 °C) temperature ranges. The Q_{10} values did not differ between elevations. Q_{10L} and Q_{10M} were negatively related only with the C:N ratio. Temperature sensitivity of decomposition of soil organic matter was not affected by bacterial activity and functional diversity (assessed using Biolog[®] ECO plates), microbial biomass or community structure (inferred from phospholipid fatty acid assays). Our findings

support a kinetics-based theory of the higher temperature sensitivity of more chemically recalcitrant soil organic matter, put forward by other authors.

Keywords Biolog[®] ECO plates · CO₂ evolution · Global warming · Microbial functional diversity · Phospholipid fatty acids (PLFA)

Introduction

Soil organic matter (SOM) decomposition processes are temperature dependent (Kirschbaum 2000). It has been suggested that the sensitivity of decomposition to global warming may be latitude-dependent (grater at higher than at lower latitude), leading to further acceleration of climate warming (Cox et al. 2000; Kirschbaum 2006). Mountain areas, with their vertical gradients of temperature and precipitation and lower variability of soil physicochemical properties, provide a platform for observing climate changes similar to those observed across latitudes, and can serve as a model for studying the temperature sensitivity of SOM along climatic (altitudinal) gradient.

Despite a number of studies on this important problem, it is still not clear which factors are the main drivers of SOM temperature sensitivity. Vanhala et al. (2007) found that old recalcitrant organic matter is more temperature-sensitive than young organic matter. Analysing the same soil by a slightly different method, Conen et al. (2008) reported no difference in temperature sensitivity between old and young SOM fractions. In our previous study of soils from an altitudinal climatic gradient in the Western Carpathians (Klimek and Niklińska 2010), we showed that temperature sensitivity of decomposition of SOM was negatively related to the soil C:N ratio, but the relationship was significant only under a level of soil moisture optimal for soil microorganisms activity. More recently, Schmidt et al. (2014) showed that the temperature sensitivity of SOM also depends on soil biology, that is, the microbial commu-

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nities. An examination of the sensitivity of the soil organic matter to changes in temperature changes should also take into account soil microbial parameters such as functional diversity or community structure.

According to van't Hoff rule, the rate of chemical processes, including metabolisms of soil microorganisms, increases two to three times with each 10 °C increase of temperature (Fang and Moncrieff 2001; Hamdi et al. 2013). However, the process rate increase is known to be higher at lower temperature ranges (Kirschbaum 1995; Niklińska et al. 1999; Tjoelker et al. 2001; Vanhala et al. 2008). It is difficult to determine exact values of this parameter, labelled Q_{10} due to differences between soils and ecosystems, and also between the mathematical models used (Bosatta and Ågren 1999; Fang and Moncrieff 2001; Tuomi et al. 2008; Karhu et al. 2010; Schipper et al. 2014). As shown by Tuomi et al. (2008), and later by Lellei-Kovács et al. (2011), a Gaussian model is the best one to use for predicting biochemical reactions to temperature, because it includes the decrease of the soil respiration rate after it reaches maximum at optimal temperature.

In our previous work (Niklińska and Klimek 2007; Klimek and Niklińska 2010) we showed that soils collected from different altitudes of a mountain altitudinal gradient did not differ in Q_{10} , against our expectation of higher temperature sensitivity in soil samples from the highest altitude studied. In those previous studies we determined Q_{10} values based on simple calculations from respiration rates at three different temperatures. We used similar methodology in these same laboratory and thus we consider the results can be compared without bias. Here we use a Gaussian model to compare the temperature sensitivity of SOM decomposition in samples from different altitudes (measured as soil respiration rate in a range of temperatures), and we analyse the relation of Q_{10} with soil physicochemical parameters and with microbial community characteristics such as microbial biomass, functional diversity and community structure.

Materials and methods

Research area and soil sampling

This work uses the data published by Klimek et al. (2015) and therefore only a brief summary of some methods is given here. The data published earlier include studied soils physicochemical properties and results of Biolog[®] ECO plates analysis. Soil samples were collected in the Beskidy Mountains, which is a part of the Western Carpathians (southern Poland). The samples were taken in forest stands on the northern slopes of five mountain massifs built of the same bedrock, known as Carpathian flysch (multi-tiered sedimentary rocks). The research sites are at approximately the same latitude, so the temperature changes associated with the latitudinal

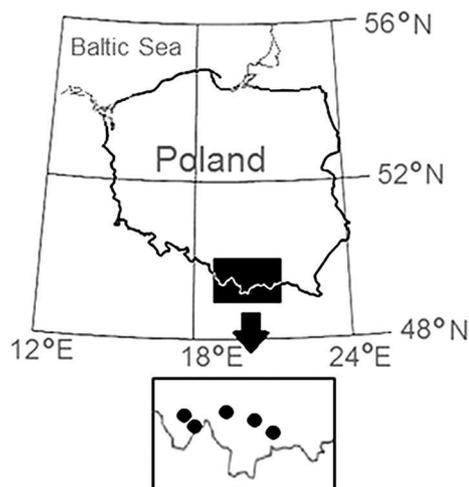


Fig. 1 Location of studied sites (mountains) in the Western Carpathians (filled dots)

gradient are negligible; we treated the five mountains as experimental replicates (Fig. 1).

Sampling plots (300 m² each) were established at three elevations (600, 900, 1200 m a.s.l.) on each mountain, giving a total 15 research sites. Mean annual temperature at those elevations is as follows: 7.2 °C at 600 m a.s.l., 5.3 °C at 900 m a.s.l. and 3.8 °C at 1200 m a.s.l. This makes a 3.4 °C decrease of mean annual temperature from the lowest to the highest stands (Klimek and Niklińska 2010). The vegetation diversity of the forest stands decreases with elevation and it changes in composition from mixed foothill forest at 600 m a.s.l. to spruce-dominated forest at 1200 m a.s.l. (Klimek et al. 2015).

At each plot, after removal of the current year's litter we collected a sample of the upper soil horizon to 10 cm depth with a steel core sampler at three points on a horizontal transect across the slope. This procedure yielded 45 separate soil samples. The data for the three soil samples per site were averaged to avoid pseudoreplication. The samples were sieved through 1-cm mesh, followed by 0.2-cm mesh, and then stored at 4 °C.

Soil physical and chemical analyses

The soil physicochemical analyses are described in detail elsewhere (Klimek et al. 2015); they included soil dry weight (DW), soil organic matter (SOM), water holding capacity (WHC), and soil pH in water and in 1 M KCl solution. Total element concentrations (Ca, K, Mg, Mn, Na and P) in each soil sample were determined using atomic absorption spectrometry or a flow injection analyser (for P content), and accuracy was verified by running five blanks and five replicates of standard certified material (CRM025-050, Sandy Loam 8, RT Corp.). Total C, N and S were analysed using a CHNS

analyser. The concentrations of dissolved organic carbon (DOC) and nitrogen (DON) were also measured.

Biolog[®] ECO plate analyses

The catabolic activity and functional diversity of the soil bacteria were analysed as described by Klimek et al. (2015). Briefly, we used Biolog[®] ECO plates containing three sets of 31 carbon substrates, with tetrazolium dye as the substrate utilisation indicator (<http://www.biolog.com>). Soil (equivalent of 5 g DW) was extracted and diluted (10^{-2}) with 0.9 % NaCl solution (pH 7), and the obtained suspension was inoculated on the plates (125 μ l per well). All plates were incubated in the dark at 22 °C, and substrate utilisation was measured as absorbance at 590 nm (μ Quant spectrometer; BIO-TEK Instruments) at 12 h intervals for 168 h. General bacterial catabolic activity was defined as sum of areas under curve for all substrates (AUC), and the functional diversity index, H'_{bact} , is based on the structure of substrate use (Klimek et al. 2015).

PLFA analysis

Soil microbial biomass and community structure were analysed based on phospholipid fatty acids (PLFA), following Frostegård et al. (1991). Total soil lipids were extracted from 0.5 g DW fresh soil using a one-phase mixture, the extracts were divided into two phases, the lower lipid-containing phase was collected, the lipid material was fractionated on columns containing silica acid, the phospholipid-containing polar lipids were collected, and a preparation of their fatty acid methyl esters was made by mild alkaline methanolysis. Methyl nonadecanoate (19:0) was the internal standard. The fatty acid methyl esters were separated using gas chromatography (Perkin Elmer, Claus 600 MS with FID detector) and quantified using qualitative fatty acid methyl esters mixes (Sigma–Aldrich).

The abundance of 22 extracted individual PLFAs was expressed as nanomoles of PLFA per DW gram soil, and their sum for each sample (PLFA_{tot}) was taken as the indicator of soil microbial biomass. Fatty acid nomenclature follows Frostegård et al. (1993). The classification of individual fatty acids to bacterial and fungal biomass ($\text{PLFA}_{\text{bact}}$ and $\text{PLFA}_{\text{fungi}}$, respectively) according to previous studies was performed according to classical studies (Frostegård et al. 1993; Hill et al. 2000; Waldrop et al. 2000). The relative shares of $\text{PLFA}_{\text{bact}}$ and $\text{PLFA}_{\text{fungi}}$ in PLFA_{tot} were also calculated.

Respiration rate measurements and Q_{10} determination

The soil samples (equivalents of 2 g DW) were placed in 50-ml glass bottles, adjusted to 50 % of their maximum

WHC and acclimated at 5, 10, 15, 20, 25 and 30 °C (± 0.5 °C) for 5 days before respiration rate measurements. Three sets of soil samples were prepared: the first set for respiration rate measurement at 5 and 10 °C, the second for measurement at 15 and 20 °C, and the third for measurement at 25 and 30 °C. These was done in order to subject the soil samples to increased temperature as briefly as possible, to prevent depletion of the available respiratory substrate (Liu et al. 2006). Carbon dioxide evolution was measured with a 30-channel MicroOxymax respirometer (Columbus Instruments, Ohio, USA). Due to the limited number of measurement channels in the respirometer, the respiration rate was measured in series with randomly selected samples. The measurements were made in darkness for 12 h to achieve stable CO_2 emission. The respiration rate of each sample was expressed as nM CO_2 per g of soil organic matter per h.

To determine the temperature sensitivity of the soil respiration rate, we fitted parameter R_0 , a and b coefficients of the Gaussian model described by Tuomi et al. (2008):

$$R(T) = R_0 e^{(aT + bT^2)} \quad (1)$$

where R_0 is the respiration rate (nM CO_2 g SOM^{-1} h $^{-1}$) at 0 °C and T is temperature Celsius. The temperature-dependent Q_{10} curves were calculated as:

$$Q_{10} = \frac{R(T + 10)}{R(T)} \quad (2)$$

based on the fitted parameters of the R(T) functions. Q_{10} values were evaluated for three temperature ranges: Q_{10L} for 0–10 °C, representing the range closest to the average annual temperature of the study area; Q_{10M} for 10–20 °C, representing the average summer temperature; and Q_{10H} for 20–30 °C, corresponding to summer heat waves in the studied area.

Statistical analyses

Differences between altitudes in mean soil physico-chemical properties and microbial characteristics (AUC, H'_{bact} , PLFA_{tot} and community structure indices, used in further statistical analyses) were compared by one-way ANOVA. One-way ANOVA also was used to test differences in parameters of the R(T) function (R_0 , a and b coefficients), respiration rates at given temperature (°C) and Q_{10} values for three temperature ranges (Q_{10L} , Q_{10M} and Q_{10H}). Right- or left-skewed data were transformed to fulfil the normality criterion. The Shapiro–Wilk test was used to confirm the normality of the data distribution within groups. The significance of differences between means was checked by Tukey's test. Differences were considered significant at $P < 0.05$.

Multiple regression analysis was used to assess the effects of soil properties on the parameters of the R(T) function (R_0 , a and b), respiration rate at 16 °C (R_{16})

and at 30 °C (R_{30}) and Q_{10} values separately. In this analysis, the R_0 , a, b and Q_{10} values were used as dependent variables. Based on multiple-variable analysis, K and DOC and the C:N ratio were independent variables representing soil nutrient status for microorganisms; H'_{bact} , bacterial biomass PLFA_{bact}, and bacterial fraction in total microbial biomass were independent variables representing community characteristics. The percentage of total variance explained by the model is reported as the R^2 value adjusted for degrees of freedom (R^2_{adj}). Statistical analyses were employed Statgraphics Centurion XVI software (StatPoint Technologies Inc., Warrenton, VA, USA).

Results

Physicochemical and microbiological properties of studied soils

Soil samples from the three altitudes differed significantly only in mean of SOM content, water holding capacity (WHC) and dissolved organic nitrogen (DON) content (Table 1). The soils from sites at the highest elevation (1200 m a.s.l.) showed the highest organic matter content (41.2 %) and higher WHC (446 %) than those at the lower two elevations. The DON concen-

tration was highest in soils at 1200 m a.s.l. (0.048 mg g⁻¹ DW), differing significantly from soils at 600 m a.s.l., which had the lowest DON value (0.021 mg g⁻¹ DW). Potassium content and soil C:N ratio did not differ between elevations (Table 1). The values for other physicochemical parameters are given in Klimek et al. (2015).

Mean H'_{bact} values differed between altitudes, ranging from 0.87 in soil under spruce forest at 1200 m a.s.l. to 1.06 in soil under mixed forest in foothills at 600 m a.s.l. (Table 1). In contrast, AUC, PLFA_{tot} and microbial community structure indices calculated from PLFA data did not differ between elevations (Table 1).

Multiple-variable analysis showed high correlations between many of the measured soil physicochemical and microbiological properties (Table S1). For example, the soil C:N ratio was positively correlated with the soil C:P and C:S ratios, PLFA_{tot} and PLFA_{fungi} and negatively correlated with soil pH. In turn, DOC was correlated positively with DON but negatively with Mg and Mn content (Table S1).

Soil respiration rate, Q_{10} and their relation to soil factors

The Gaussian model of the temperature dependence of soil respiration fit the measured data well, and the models fitting adjusted for degrees of freedom (R^2_{adj})

Table 1 Climatic data (mean annual temperature and precipitation) and means (\pm SD) for selected physicochemical and microbial indices as well as temperature sensitivity of organic matter in soils from stands along an altitudinal gradient in the Western Carpathians (n = 5)

	P value	Altitude (m a.s.l.)					
		600	900	1200			
Climatic property and its unit							
Mean annual temperature (°C)	–	7.2	5.5	3.8			
Mean annual precipitation (mm)	–	1000	1210	1420			
Soil property and its unit							
Physicochemical							
SOM (%)	0.0141	26.2 a	(8.4)	23.0 a	(5.1)	41.2 b	(11.4)
WHC (%)	0.0004	177 a	(48)	218 a	(54)	446 b	(121)
DOC (mg g ⁻¹ DW)	0.3744	0.31	(0.08)	0.36	(0.08)	0.39	(0.11)
DON (mg g ⁻¹ DW)	0.0367	0.021 a	(0.005)	0.035 ab	(0.014)	0.048 b	(0.026)
K (g kg ⁻¹ DW)	0.2802	0.36	(0.15)	0.46	(0.17)	0.30	(0.05)
C:N	0.2422	18.4	(2.8)	17.4	(2.7)	20.1	(1.7)
Microbiological							
AUC	0.0721	41.5	(15.9)	39.0	(16.6)	19.4	(11.8)
H'_{mic}	0.0219	1.06 b	(0.16)	1.07 ab	(0.08)	0.87 a	(0.09)
PLFA _{tot} (nM g ⁻¹ DW)	0.7199	29.1	(6.6)	26.8	(6.2)	29.7	(4.3)
PLFA _{bac} (nM g ⁻¹ DW)	0.6345	15.2	(2.4)	15.5	(4.5)	17.0	(1.8)
PLFA _{fungi} (nM g ⁻¹ DW)	0.4704	10.9	(3.6)	8.8	(1.7)	10.0	(2.5)
Bacteria share (%)	0.1208	52.7	(4.7)	57.4	(3.8)	57.5	(2.6)
Fungi share (%)	0.2156	36.9	(4.3)	32.9	(3.2)	33.5	(3.4)
Temperature sensitivity							
R_0 (nM CO ₂ g ⁻¹ OM h ⁻¹)	0.1067	0.34	(0.33)	0.29	(1.18)	0.86	(0.64)
R_{16} (nM CO ₂ g ⁻¹ OM h ⁻¹)	0.0497	1.48 a	(1.11)	1.32 a	(0.64)	3.69 b	(2.27)
R_{30} (nM CO ₂ g ⁻¹ OM h ⁻¹)	0.0373	3.82 a	(2.85)	3.02 a	(1.26)	9.65 b	(5.92)
a	0.6496	0.123	(0.029)	0.121	(0.021)	0.110	(0.021)
b	0.5894	–0.0012	(0.0006)	–0.0012	(0.0005)	–0.0009	(0.0005)
Q_{10L}	0.5584	3.21	(0.82)	3.08	(0.51)	2.79	(0.44)
Q_{10M}	0.5928	2.46	(0.31)	2.35	(0.21)	2.31	(0.17)
Q_{10H}	0.5677	1.93	(0.10)	1.84	(0.15)	1.94	(0.20)

Values bearing different bolded letters differ significantly (ANOVA with Tukey's HSD test)

explained from 97.0 to 99.6 % variability in given temperatures (data not shown). R_0 representing the respiration rate at 0 °C, serving as a proxy of SOM decomposability, did not differ between elevations ($P = 0.1067$) (Table 1). At higher temperatures (≥ 16 °C) there were differences in mean modelled respiration rate between soils from different altitudes (Table 1). In the 16–30 °C range the modelled soil respiration rate was significantly higher for soils from 1200 m a.s.l. than for soils from lower altitudes (Fig. 2). The respiration rate of soils from 1200 m a.s.l. was about three times higher than that of soils from lower altitudes (Table 1).

Coefficients a and b of the Gaussian model did not differ between elevations (Table 1). For soil from all elevations, the temperature-dependent Q_{10} values calculated based on the respiration rate values fitted to Gaussian functions showed decreasing temperature sensitivity with the increase of test temperature (Fig. 2). The differences in Q_{10} between altitudes were not significant in any of the tested temperature ranges (Table 1).

Multiple regression yielded significant models for R_0 , R_{16} , and R_{30} and for Q_{10L} and Q_{10M} values (P values ranging from 0.0384 to less than 0.0001). The models explained from 25.5 % (Q_{10L}) to 87.4 % (R_0) of the variance of these values (Table 2). R_0 was correlated positively to the DOC content and the soil C:N ratio (as well as R_{16} and R_{30} were; data not shown). Q_{10L} and

Q_{10M} were correlated only (and negatively) with the soil C:N ratio (Table 2).

Discussion

The soils from the studied altitudes differed significantly in only some of the measured physicochemical and microbiological properties. Although the respiration rate found for soil from the highest studied altitude (1200 m a.s.l.) was higher than for soil taken from lower altitudes, we did not find differences in temperature sensitivity of decomposition of SOM between elevations; the study did not confirm higher temperature sensitivity of soils from colder sites than from warmer sites. This result is in line with findings from our earlier study of an altitudinal gradient in the Western Carpathians (Niklińska and Klimek 2007; Klimek and Niklińska 2010).

In a current study, the calculated Q_{10} values were temperature-dependent, decreasing with the increase of test temperature. In the 0–10 °C range the Q_{10} values were 2.8–3.2, exceeding the 2.5 value reported as a mean for forest ecosystems (Hamdi et al. 2013) and the 2.0–2.5 range often used in global C cycling models (Ise and Moorcroft 2006). Although the difference in Q_{10} is not large and the soil respiration rate at lower temperature is generally low, such temperature sensitivity may herald a pronounced increase in soil carbon loss to atmosphere (Kähkönen et al. 2001), because such conditions are typical for the large earth areas during the greater part of the vegetation season.

We expected that temperature sensitivity of soil organic matter decomposition processes along mountain altitudinal gradient would differ between altitudes. The lack of differences in Q_{10} between altitudes, as we found, may be due to too short studied gradient but also differences between stands from the same elevation. On the other hand, the lack of these differences may result from stands unifying, resulting from human activity (e.g. selective tree cutting or spruce introduction below natural belt) that altered the soil processes of some stands. Such anthropoppression may be reflected in the lack of large differences in soil physicochemical and microbiological characteristics between altitudes.

Multiple regression indicated that the soil respiration rate modelled with a Gaussian function was positively correlated with the soil C:N ratio and DOC concentration. This demonstrates the importance of readily degradable carbon compounds for soil microorganisms metabolic activity. Demoling et al. (2007) found that soil bacteria can be limited by carbon availability even in soils with a wide C:N ratio. In our current study on soils from mountain altitudinal gradient, the soil C:N ratio was positively correlated with soil microbial biomass, including fungal biomass as measured by the PLFA method. High microbial biomass is related to high microbial activity, but carbon availability may affect microbial activity and biomass differently (Allen and

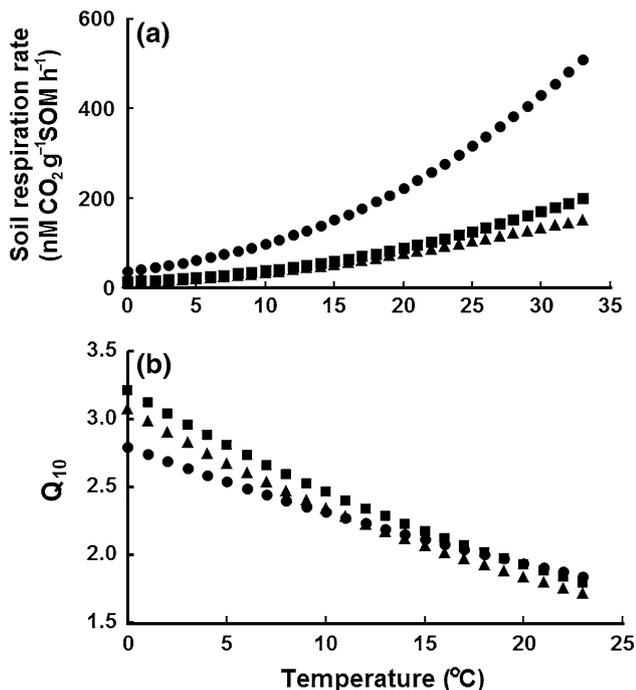


Fig. 2 Relation of **a** soil respiration rate and **b** soil temperature sensitivity to temperature in soils from 600 (filled squares), 900 (filled triangles) and 1200 m a.s.l. (filled dots). Gaussian functions describe temperature dependence of soil respiration. The Q_{10} curves were calculated using R_0 , a and b parameters of Gaussian function averaged ($n = 5$) for each of the studied elevations

Table 2 Relationships between R_0 and Q_{10} in low and medium tested temperature ranges (Q_{10L} , Q_{10M})—multiple regression results (only the significant variables are presented)

Parameter	Equation parameter	β	<i>P</i> value	(R_{adj}^2) (%)	P_{mod}
R_0	Intercept	-1.279	0.0037	87.4	< 0.0001
	C:N	1.851	0.0001		
	DOC	0.091	0.0001		
Q_{10L}	Intercept	5.358	0.0002	23.5	0.0384
	C:N	-0.125	0.0384		
Q_{10M}	Intercept	3.374	< 0.0001	30.5	0.0191
	C:N	-0.054	0.0191		

The regression coefficients (β) and levels of significance for the model (p_{mod}) and for particular model parameters (p) are also given

Schlesinger 2004; Spohn and Chodak 2015). The soils from the highest altitude studied, having higher SOM content, showed a higher respiration rate modelled with a Gaussian function for temperatures above 16 °C. However, the soils from the different altitudes did not differ significantly in C:N ratio, DOC content or microbial biomass. These results suggest site-specific differences unrelated to their substrate (the sites are all on the same bedrock, Carpathian flysch).

Multiple regression analysis indicated that Q_{10} in the low and medium temperature range was correlated negatively with the soil C:N ratio. Such a trend, opposite to the results for respiration rate, seems to support the kinetic-based suggestion of higher temperature sensitivity of more chemically recalcitrant SOM and its associated soil microbial communities (Vanhala et al. 2007). That idea was reconsidered recently by Xu et al. (2015), who made a long-term laboratory study of the temperature sensitivity of soils along an altitudinal gradient in the Wuyi Mountains (China). They showed that soils with higher soil organic carbon recalcitrance showed significantly higher temperature sensitivity of decomposition of SOM. Their results also indicate that the temperature sensitivity of both the labile and the recalcitrant SOM fractions is negatively correlated with the soil C:N ratio in stands delineated at these same altitudes. Similar results were given by Hiltunen et al. (2013) in their study of temperature sensitivity along the peat profile; the oxic peat layer, with the lowest C:N ratio, showed the highest temperature sensitivity. A low soil C:N ratio may, however, have two different explanations driven by completely different ecological processes. One explanation is that a high rate of SOM decomposition causes depletion of thee carbon and a relative increase of nitrogen; this can take place in conditions favourable for fast decomposition. A low soil C:N ratio may also result from the chemical composition of litter, and may depend more on the quality of litter produced in a given ecosystem.

Against expectation, we did not find that the Q_{10} coefficients were related to bacterial functional diversity. Apparently, more functionally diverse communities are able to effectively degrade various organic substrates and, as a result, can take more advantage of increasing

temperature. The lack of such relationships may result from carbon availability restrictions influencing bacterial performance, or may be due to limitations of the method, that is, the Biolog[®] plates (Campbell et al. 2003). Vanhala et al. (2008) also found that the results from Biolog[®] plates did not correlate with Q_{10} based on a Gaussian function. Our findings may also suggest a stronger effect on temperature sensitivity of decomposition of soil organic matter from the fungal part of community than from bacteria. There is in line with findings of Pietikäinen et al. (2005), who found that reaction of growth rate to temperature increase was more drastic for fungi than for bacteria. That possibility needs to be studied further.

Conclusions

The Q_{10} parameters describing temperature sensitivity of SOM were temperature-dependent and decreased with the increase of test temperature. In the lower temperature range, typical for the study area, we found the Q_{10} values of around 3, higher than the values typically used in climate models. We did not find differences in Q_{10} between altitudes differing in mean annual temperature, which suggest a lack of higher temperature sensitivity in soil of colder regions or may result from too short climatic gradient studied. The Q_{10} values in the low and medium temperature ranges (0–10 °C, 10–20 °C) were negatively correlated with the C:N ratio. These findings are in line with the kinetics-based suggestion of the higher temperature sensitivity of more chemically recalcitrant SOM and its associated soil microbial communities.

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